

Dust aerosol properties on Mars from spacecraft: A current perspective

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- Context
- History
- The planet Mars
- Radiative transfer & aerosol properties
- Retrieving aerosol properties from spacecraft
- Our current understanding







Why studying dust is important?

Airborne dust absorbs and scatters light \Rightarrow Local warming & cooling in the atmosphere

It modulates thermal and dynamical structure of the atmosphere

 \Rightarrow affects atmosphere and climate



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History (1/2)

Mariner 4

(1965)

- First flyby: 22 pictures
- First detailed picture of the surface
- \Rightarrow dull & cratered planet (Moon-like appearance)

Mariner 6 & 7 (1969)

- Other flyby pictures: same dull terrains
- But thermal profiles suggest the presence of c
 airborne aerosols







History (2/2)

- Mariner 9 (1971-72)
 - First orbiter, arrived during a global dust storm (GDS) event
 - About 20,000 pictures, UV & IR spectroscopy
 - Observed the global dust storm (GDS) decay, dust strongly influences temperature
- Vikings 1 & 2 (from 1976 to early 80's)
 - 2 orbiters & 2 landers
 - observed GDS events, study impact on atmosphere & climate
- Start of continuous monitoring (from 1997)
 - started with Pathfinder & Mars Global Surveyor
 - since many more (Odyssey, MER, MRO, MEx, Phoenix, MSL, MOM, MAVEN, TGO, Insight, ...)



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The planet Mars



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Mars



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Mars in the Solar system

4th planet of the Solar system

• Telluric planet

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- Diameter: 3394 km 0.53 x Earth
- Masse: 6.2 x 10²³ kg 0.11 x Earth
- Gravity: 3.72 m/s² 0.38 x Earth







Orbit and rotation

Revolution around the Sun





Rotation on axis

1 Martian day or sol: ~ 24h40 min





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Asymmetry between North & South hemispheres

+ 45% more solar energy received at perihelion than at aphelion

- South: Extreme seasons
 - North: temperate seasons

milder winter and cooler summer

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cold winter and warm summer







Progression in the Martian year

- \Rightarrow Solar longitude (L_s)
 - 1 Martian year = 360° of L_s





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Martian dust

Dust is ubiquitous on Mars !



Lift in the atmosphere by winds and storms!

t - ۵۶ Dust devils by MER Spirit Credit NASA/JPL

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Non dusty season

Occurs during the "colder" season around Aphelion



Period characterized by:

- Low dust loading
- Repeatable from year to year (*)
- Local storms and dust devils (*varying patterns)





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Dusty season & dust storms

Dusty season: occurs during

the "warm" season around Perihelion



- Higher dust loading
- Important interannual variability
- Dust storm spatial extent
 - Local (frequent)
 - Regional (usually 1-3 per year)
 - Global (not every year, 1 out of 3 year in average)
- Storms often form at the edge of the polar caps



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Radiative transfer









Radiative transfer

Study the interactions between the light and the atmosphere & surface

 \Rightarrow Different types of interaction: absorption, scattering, reflection (scattering from a surface)





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Light-dust interaction

Light is an electro-magnetic wave defined by its wavelength





Each molecule/particle has a specific interaction with light that is wavelength dependent & related to its intrinsic properties:

- Composition
- Size

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• Shape



Dust properties (1/3)

From Wolff et al. 2009

Composition:

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- Real refractive index: light bending (refraction)
- Complex refractive index: absorption
- \Rightarrow single scattering albedo (SSA): scattered vs absorbed fraction





Dust properties (2/3)

Particle size

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- extinction cross-section: ability to interact with light
- Size parameter: $X = \frac{2\pi r}{\lambda}$

Martian dust particle size:

- $r_{eff} = 1-2 \ \mu m$ typically
- smaller particles at higher altitudes
- larger particles up to 4 μm during storms near surface

Distributions: $r_{eff} \& v_{eff}$ (1st & 2nd moms) as defined in [Hansen and Travis, 1974]



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(r: particle radius; λ : light wavelength)





Distributions: r_{eff} & v_{eff} as defined in [Hansen and Travis, 1974]



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Dust properties (3/3)

Particle shape

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• Phase function: angular probability distribution for scattering

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Influence of the shape on the phase function

Phase function comparison of cylindrical prisms with a sphere

using cylindrical prisms with different D/L ratio (diameter/length) [Wolff et al., 2009, 2010]







Dust properties from spacecraft









Sparse measurements of orbiters and combination with landers allowed to:

- Observe a global dust storm (GDS)
 ⇒ importance of dust & radiative properties
- Deduce the general seasonal & spatial trends of dust cycle
- Estimate the particle size: r_{eff} =1.6-1.8 µm & v_{eff} =0.2-0.6

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Beginning of global coverage

Using MGS Mars Observer Camera (MOC, visible imager) and Thermal Emission Spectrometer (TES, IR spectrometer)

- Possible to monitor globally the dust activity (from MOC)
 ⇒ Dust storms are occurring constantly (not only during "storm season")
- Daily quantitative retrievals of dust opacity column by TES

Nadir viewing

- Along with temperature (T) (and water ice)
- \Rightarrow Seasonal dust cycle & spatial distribution



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Dust cycle from TES (& other cycles)

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Non dusty season: annual repeatability Dusty season: variable

Dust & T correlation

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Dust & ice clouds anticorrelation



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EPF (Emission Phase Function) measurements were performed by TES

Aiming the same location with different angles

⇒ Study dust properties (PF, size, shape)



- Solar band channel (0.3-3 μm)
 ⇒ distinction of aerosol types (dust/ice cloud)
 - Mean size: r_{eff}=1.6 μm & v_{eff}=0.2-0.4
- Dust storm size: larger particles 2-3 μm
- Shape: cylinder particle (moderate axial ratio)

From [Clancy & Wolff, 2003]



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EPF viewing

© Clancy and Lee, 199



From [Wolff & Clancy, 2003]



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Dust properties from TES (3/4)

Distinction between dust and ice clouds



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Mars Aerosol Type & Optical Depth (TES EPF analysis, 1999-2001)

Ice cloud (1-2 μm) Ice cloud (3-4 μm) Dust (~1.6 μm) Bright dust (1.0-1.2 μm)



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Altitude distributions

Mars-Express (Mex) (since 2003) SPICAM UV-NIR spectrometer

Occultation

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- Observing sunset/sunrize through the atmophere
- Direct attenuation of the solar/stellar radiation
- \Rightarrow No RT calculation! (=>no complications)

⇒ Altitude distributions (opacity, size)









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Altitude distributions

Mars Reconnaissance Orbiter (MRO) (since 2006) MCS: IR radiometry

Limb

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- Observing the limb of the planet
- Emission by dust & atmosphere (FIR)
- \Rightarrow Require RT calculation! (=> complications)

\Rightarrow Altitude distributions





Related to work of [Kleinböhl et al., 2009]





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Refining dust properties

Using MRO/CRISM visible-NIR spectroscopy

- EPF measurements
- worked during dust storm
- combined with ground truth opacities from MERs
- \Rightarrow complex refractive index & SSA
- \Rightarrow size (1.8 µm average for the observation used) & shape (cylinder with moderate axial ratio)





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Improved dust properties in modeling

Global circulation models (GCM) are used to simulate and understand the chemical and physical processes occurring in the Martian atmosphere (see Lori's presentation)

- Use of the improved dust properties
- ⇒ helped to solve the temperature anomaly observed between T_{TES} & T predicted by models



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Work of Madeleine et al., 2011



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RoadMap project

Perform laboratory measurements on Martian dust analogues to improve retrieval and modeling:

- Lab measurements include:
 - Radiative & scattering properties (SSA, phase function, polarization)
 - Dust lifting characterisation, deposition, dynamics
 - Micro-physical properties
- Dust related goals of "atmopheric measurements Work Package"
 - Test the impact of new properties on nadir retrieval
 - Deduce particle size & aerosol's composition from occultations
- Using TGO/NOMAD (from 2018)
 - UV-visible and NIR spectrometer





Aerosol altitude distribution from NOMAD

The 10,000+ TGO/NOMAD-UVIS occultations are used to produce a complete climatology of the aerosol altitude distribution.

No dust/ice cloud distinction!

 \Rightarrow Distinction possible combining with NOMAD-SO

Ongoing work...

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Work of Flimon et al., in prep





Particle size is also deduced from TGO/NOMAD-UVIS occultations

Combining with NOMAD/SO will also improve the particle size retrieval and allow to detect larger particles



Assimilated dust climatologies

The work of [Montabone et al. 2020] combines the retrieval of the several instruments to be assimilated in order to produce a complete seasonal climatology of the dust opacity on more than 13 MYs





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Dust column from EMIRS

EMM/EMIRS TIR spectrometer

- Dust OD at 9.3 µm
- Provide global synoptic view at all local times (day & night)

Allowed to observe the initiation, growth, and decay of the early regional dust storm of MY 36



Dust Column Optical Depth (1075 cm⁻¹)

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Dust properties from lab measurements

In the frame of roadmap, laboratory experiments were performed on 3 Martian dust analogues to measure the:

- Phase functions & polarization
- Complex refractive index
 Then combined with realistic particle shapes
 => New set of scattering properties were calculated

Comparing results:

- Analogues don't seem suitable for UV range were calculated (SSA>0.9, appear too bright)
- Better agreement for visible





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Work from Martikainen et al., 2022



Improving dust properties

Current work assesses the impact of using more realistic "hexahydra" particle shapes in comparison with cylinder shapes of [Wolff et al. 2009]

\Rightarrow Absorption

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Primary impact <0.4 μ m & 2.9 μ m band (new shapes with many edges induce additional scattering => require more absorption)

\Rightarrow Phase function

- Very limited impact for phase angles <150°
- Relatively limited at larger angles



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- What physical processes control the initiation, growth, and decay of dust storms?
- What physical processes control the interannual variability of global dust storms?
- What is the current global dust budget and how has it evolved over time?
- How do dust particle size distributions evolve over time in different conditions (storms, dust devil injection, etc.)?





Some resources to consider

• Dust climatology of [Montabone et al., 2015,2020] https://www-mars.lmd.jussieu.fr/mars/dust_climatology/

 Montabone, L., Forget, F., Millour, E., Wilson, R.J., Lewis, S.R., Cantor, B., Kass, D., Kleinböhl, A., Lemmon, M.T., Smith, M.D., Wolff, M.J., *Eight-year Climatology of Dust Optical Depth on Mars.*

Icarus 251, pp. 65-95 (2015), doi: https://doi.org/10.1016/j.icarus.2014.12.034

Montabone, L., Spiga, A., Kass, D. M., Kleinböhl, A., Forget, F., Millour, E.,
 Martian Year 34 Column Dust Climatology from Mars Climate Sounder Observations: Reconstructed Maps and Model Simulations.
 J. Geophys. Res. - Planets (2020), doi: https://doi.org/10.1029/2019JE006111

Hexahydrae shape database https://zenodo.org/record/4711247#.YhYXbS1Q3RU

Saito, M., P. Yang, J. Ding, and X. Liu, A comprehensive database of the optical properties of irregular aerosol particles for radiative transfer simulations. J. Atmos. Sci., in press, doi: https://doi.org/10.1175/JAS-D-20-0338.1

- DISORT radiative transfer software http://www.rtatmocn.com/disort/ + python wrappers (https://github.com/kconnour/pyRT_DISORT)
- "Scattering by Particles & Surfaces" softwares of M. Mishchenko https://www.giss.nasa.gov/staff/mmishchenko/brf/



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